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Introduction

Figure 1: Topographic map of the upper Blue Nile basin

38° E

39° E

40° E



Elevation (m)





-12° N		
-11° N		
-10° N		
-9° N		
-8° N		





The evaporation is generally underestimated over the upper Blue Nile basin. The basin's annual water budget can not be balanced using these datasets with an error of 30% of the mean annual precipitation. **UBN** Evaporation 120 **WM Evaporation Data NTSG Evaporation Data** 100 onth) 80 (mm/me **60**₿ ☞----Depth 40 Alexi Evap ——— NTSG Evap — — — → RegCM3 Evap 20 WMEvap Μ Μ 0 D S Month







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Introduction



UBN Potential Evaporation

Figure 2: Comparison between the temporal and spatial distribution of several global evaporation datasets







In this analysis we use the following data: (i)TRMM v7 Multi-satellite Precipitation Analysis (TMPA) 0.25° × 0.25° resolution 3B42 (Huffman et al., 2007) (ii)CRU TS 3.1 potential evaporation dataset (Mitchel and Jones, 2005) (iii)GRACE Terrestrial water storage (Chambers, 2006) (iv)WM Evapotranspiration data (Willmot and Matsuura, 2011) (v)Cropland and Pasture data 1700-2007 (Ramankutty and Foley, 1999) (vi)Dominant Flow Routing (DRT) algorithm (Wu et al., 2011) (vii)HWSD Water holding capacity (FAO, IIASA, ISRIC, ISSCAS & JRC, 2009) (viii)NASA-SRB surface shortwave and longwave radiation (Darnell et al., 1996; Gupta et al., 1999)

An optimization model is formulated to minimize the weighted mean-squared deviation of the estimated hydrological variables from the input data for a typical year over the upper Blue Nile Basin.



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Data

Methodology

: pixel basin storage misfit

- The monthly pixel water balan
$DS_{n,m} = P_{n,m} + Q_{in_{n,m}} - ET_{n,m}$
- The tributary flow constraint:
$Q_{in_{n.m}} = \sum \left(\Delta t \times \left(P_{n,m} - ET_{n,m} \right) \right)$
<i>trib</i> - The basin outflow flow constr
$R_{m,g} = \sum Q_{out_{g.m}}$
- The Evapetranspiration const
$ET_{n,m} = ET_{crop_{n,m}} + ET_{noncr}$
- The Radiation constraint:
$ET_{n,m} \times \lambda \leq Rad_{net_{n,m}}$
 The storage threshold and no
- The storage threshold and no $\Delta S_{\min} \leq \Delta S_{n,m} \leq \Delta S_{\max}$
- The storage threshold and no $\Delta S_{\min} \leq \Delta S_{n,m} \leq \Delta S_{\max}$ $S_{n,m} \leq S_{threshold}$
- The storage threshold and not $\Delta S_{\min} \leq \Delta S_{n,m} \leq \Delta S_{\max}$ $S_{n,m} \leq S_{threshold}$ $S_{n,m}, P_{n,m}, R_{m,g}, Q_{in_{n,m}},$

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Physical and Hydrological constraints

ce:

 $-\Delta S_{n,m}$, n,m

raint:

traint:

 $_{m} + ET_{Lake_{n.m}} \& ET_{crop_{n.m}} = K_{crop_{n.m}} \operatorname{PET}_{n,m} \operatorname{\frac{A}{Crop_{n.m}}}$ $rop_{n.m}$

on-negativity constraints :

 $Q_{out_{n.m}}, PET_{n,m}, ET_{n,m}, ET_{crop_{n.m}}, ET_{noncrop_{n.m}}, ET_{Lake_{n.m}} \ge 0$

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Physical and Hydrological constraints

Several flow stations were added to the optimization model to enhance the model flow routing. The basin is divided into 5 sub-basins to investigate the spatial and temporal evaporation trend variation between the 5 sub-basins.

Figure 3: Flow gauge stations locations and the sub-basins created

A comparison between the spatially averaged input and assimilated hydrological variables is shown in Figure 4. The model finds that TRMM overestimates precipitation by 9%. It was found that the model annual evaporation estimate agrees with the ALEXI evaporation product (Anderson et al., 1997). However the spatial distribution

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Model Results

UBN Evaporation

Figure 5: A comparison between the spatial distribution of the model and ALEXI product annual ET

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Model Results

Model Annual ET

A PATA 500 60 800 950 100 250 100 DATA 500 600 800 950 100 250 100

Model Annual ET

Civil & Environmental Engineering

The climatology of the sub-basins is quite different, the western basins are generally wetter than the eastern highlands. The evaporation over the BK and KK sub-basins has a higher seasonal pattern where it peaks around the rainy season while the peak evaporation extends over 5 months in the MD and KM. KK Evaporation **BK Evaporation**

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Model Results

higher than the current upper Blue Nile croplands.

Depth (mm)	
Precipitation	
Evapotranspiration	
Runoff	

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Model Results and Discussion

The model results highlight the role played by natural vegetation in the basin's water budget. Natural vegetation generally evaporates

Figure 7: MODIS 2009 Crop land use

Summary of Results

Table 1. Comparison between the annual water budget depths over the UBN

Data	Model
1362	1429
656	952
276	278

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Table 2. Summary of the annual water budget depths for 5 sub-basins

Depth (mm)	BD_Kes	Kes_Kar	Kar-Man	Man_Die
Precipitation	1076	1188	1338	1302
Evapotranspira	829	979	1085	939
Runoff	269	202	251	357
RC %	25	17	19	27

1. The available global satellite evaporation datasets generally underestimate the evaporation from the upper Blue Nile basin.

Conway, D., 1997, A water balance model of the Upper Blue Nile in Ethiopia, Hydrological Sciences Journal, 42(2), 265-286 Conway, D. (2000), The Climate and Hydrology of the Upper Blue Nile River. The Geographical Journal, 166: 49–62. McLaughlin, D. (1995), Recent developments in hydrologic data assimilation, Rev. Geophys., 33(S2), 977–984 Ramankutty, N., and J. A. Foley (1998), Characterizing patterns of global land use: An analysis of global croplands data, Global Biogeochem. Cycles, 12(4), 667–685 Ramankutty, N., and J. A. Foley (1999), Estimating historical changes in global land cover: Croplands from 1700 to 1992, Global Biogeochem. Cycles, 13(4), 997–1027 Tafesse, T., 2001, The Hydropolitical Assessment of the Nile Question: An Ethiopian Perspective, Water International, 26(4), 1-11

Summary of Results

Conclusions

2. Natural vegetation plays the most important role in the hydrological budget of the upper Blue Nile basin.

3. The current croplands and cropping patterns are more water efficient than the natural vegetation in the upper Blue Nile basin

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